

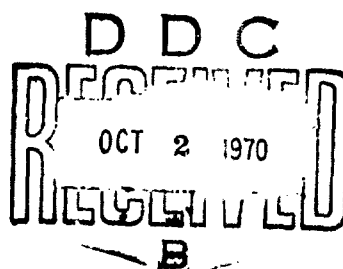
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Special Report 28

JULY, 1954

**Measurements of
Ice Tunnel Deformation
Camp Red Rock,
Greenland**

by Robert E. Hilty



**U. S. ARMY SNOW ICE AND PERMAFROST
RESEARCH ESTABLISHMENT**

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PREFACE

This is a report of work performed during the 1957 Greenland field season on USA SIPRE Project 22.1-24 and CE Project 24, Ice cliff studies. The purpose of these investigations was to obtain basic information regarding ice motion, recent history of the ice edge, and the physical characteristics of the ice, as a basis for further engineering research. Part of the work was performed under Contract DA-11-190-ENG-19 with Ohio State University.

The research work was performed by Mr. Robert E. Hilty.* Work on this project was performed for the Snow and Ice Basic Research Branch, Mr. James A. Bender, chief.



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ICE TUNNEL DEFORMATION MEASUREMENTS
CAMP RED ROCK GREENLAND

by
Robert B. Hilly

INTRODUCTION

Detailed scientific studies on the regimen of the cliff terminus of the North Ice Cap in Greenland were initiated as a full scale project (CE Project 24) in the summer of 1955, at Camp Red Rock, Nunataarsuaq, 40 miles northeast of Thule Air Base. One of the several means of gathering information was through deformation studies in a 30-m deep horizontal tunnel excavated normal to the trend of the cliff face and aligned essentially parallel to the direction of glacier flow. During the spring and summer of 1956, a tunnel addition (referred to here as the 1956 tunnel) was extended from the side of the 1955 tunnel. The excavation in 1956 was dug to a lower level and exposed approximately 15 m² of bouldery subglacial floor. This provided an excellent opportunity to obtain precise and detailed information on basal glacier movement as well as to study the effects on a patch of frozen ground suddenly relieved of high stresses. The tunnel walls and ceiling were instrumented with several rings of motion pegs which were surveyed with a theodolite from a system of stationary stakes. Mapping of the ice stratigraphy and other deformation studies augmented the motion survey.

The primary objective of Project 24 in 1957 was to resurvey the position of the pegs in the 1956 tunnel. Compilation of data and conclusions derived constitute the basis of this report.

EXCAVATION

The 1956 section of tunnel was begun in order to expose the downward traces of heavily dirt-laden shear planes, prominent in the 10-m area of the 1955 tunnel, and to discover their configuration and the nature of bottom contact (Fig. 1 - 2).



Figure 1. Looking south in tunnel,
August 1956.



Figure 2. Looking south in tunnel, June 1957.
Notice bowed-in west wall and boulder floor.

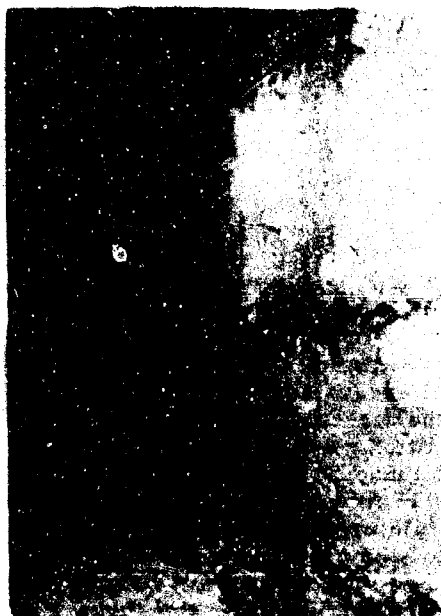


Figure 3. Dirt layers in west wall near ring 4.

The extension started perpendicular from the east wall at about 9 m in from the mouth of the 1955 tunnel. The shaft gradually curved to a northerly direction and descended steeply in an attempt to remain in the plane of the waning dirty ice bands. The principal portion of the 1956 excavation extended 15 m northward from the descending curve and was orientated essentially parallel to the 1955 tunnel (Fig. 6). The rocky floor of the tunnel sloped slightly northward about 70 cm over the 15-m length. The height of the tunnel decreased from 5 m above the floor at the south end to 2 m above it at the north end.



Figure 4. "Suspended" dirt layers on west wall near ring 3. Note sand and ice floor on left and boulder floor on right.

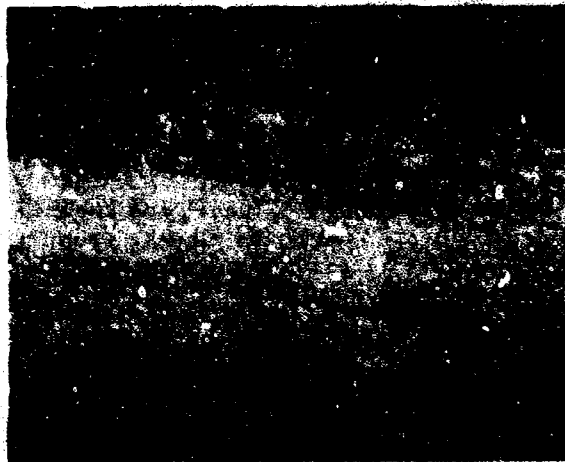


Figure 5. Ice stratigraphy and interbedded dirt bands, on west wall 2 m south of ring 4.

STRATIGRAPHY

The shear planes of the 1955 excavation terminated abruptly, shortly below the 1955 tunnel level. However, another set of dirt layers, not connected to those mentioned above, were interbedded in the ice at the south end of the tunnel (Fig. 3-5, 7). These dirt bands consisted of solid pebbly sand ranging from a few centimeters in thickness throughout most of the layer to a maximum of 20 cm in the irregularly spaced pods. The dirt layers dipped northward in the same direction as the ice foliation. Ice-dirt contact was always sharp and conformable to slightly non-conformable. No folding was apparent in the dirt bands. The ice varied from green to white to gray and dipped approximately 7°N at the north end of the tunnel and up to 23°N at the south end nearer the ice terminus. Original dips taken on the dirt bands and ice foliation were remeasured in 1957 and no appreciable differences could be detected.

Lichen-covered boulders up to 1 m in diameter constituted most of the tunnel floor (Fig. 2). A smaller area in the southern portion was covered with a thick blanket of interbedded sand and ice layers (Fig. 4).

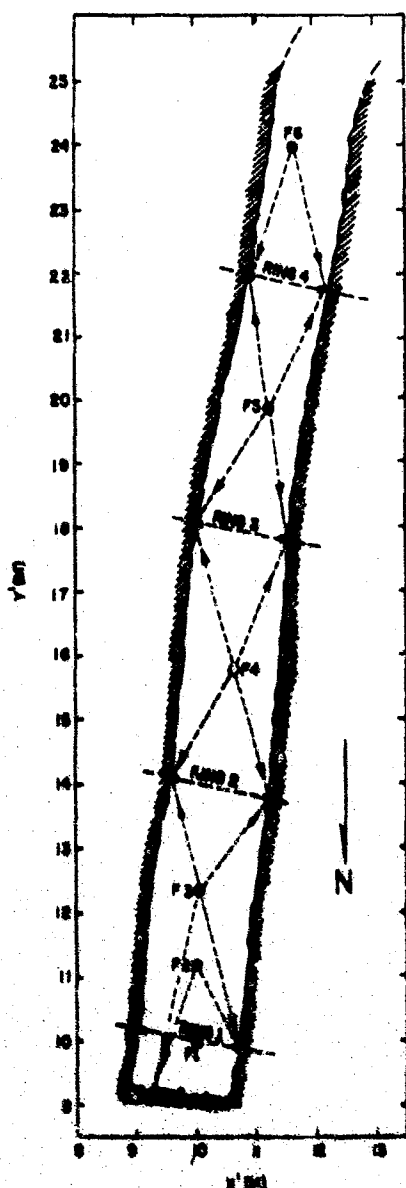


Figure 6. Diagram of 1956 tunnel showing rings of pegs and triangulation network in relation to coordinate system.

It is postulated that the ice and dirt sequence exposed on the walls of the 1956 tunnel formed from a snow drift overridden by and incorporated in the advancing North Ice Cap. Facts supporting this hypothesis are: a) the highly folded dirt bands in the 1955 tunnel did not reach the basal part of the glacier; b) the dirt bands in the 1956 tunnel were unfolded, had a low dip, and appeared to be depositional rather than tectonic features; c) these suspended dirt layers appeared to have had no former connection with the subglacier terrain. In a plunge pool 140 m west of the tunnel and at the base of the cliff, a very similar stratigraphic section was observed in a snow drift which extended directly under the ice cliff.

Not only did the tunnel cut provide an excellent exposure of basal ice stratigraphy, but the stable floor furnished the possibility of an internal baseline for precise deformation studies under approximately 40 m of ice and under uniform temperature conditions on the order of -10°C .

INSTRUMENTATION

Four rings of pegs were set in the walls at 4-m intervals; crossed rows of pegs were lined along the back face, and auxiliary pegs were placed at the south end of the west wall (Fig. 7). Because of lack of time, only a majority of the ring pegs and the back face pegs were resurveyed. Plumb and tape measurements were taken on the auxiliary pegs. Absolute motion was triangulated on 35 pegs in mid-August 1956 and again in early June 1957. The time interval amounted to an average of 294 days.

Pegs consisted of 15 cm long by 1 cm diam wooden dowels soaked in water and forced into a hole of the same diameter. Steel pins driven into the ends of the dowels were clipped to a point to provide a sharp target. The floor pegs employed as triangulation stations were 2.5 cm diam steel rods driven securely between the boulders.

A Wild T-2 theodolite was used to triangulate each peg from two floor stations (Fig. 6). The single set of direct and reverse angles taken were within 30 sec difference with very few exceptions and usually within 10 sec difference. Considering this amount of variation over the short distances used, the peg coordinates presented are accurate to the nearest centimeter. Rough plumb and tape measurements on the west wall of rings 1 and 4 check within 5% of those obtained by triangulation.

The only case where the triangulated peg positions are beyond reasonable accuracy is on the back face. Angles from both stations were in the range of 180° and the resulting discrepancies in the peg coordinates render them unreliable.



It is evident that the floor did not respond to the removal of the ice load. Taping and triangulation of the floor stations for a second time in 1957 reveal no differences. The close measurements on the "A" pegs also indicate lack of floor heaving since it would in turn affect the angles taken. Visual observations showed no shifting of boulders.

Coordinates of each peg are presented in Table I. The northernmost floor station (F1) was arbitrarily chosen as the reference point and was given a value of 10 m for X', Y', and Z' directions. The coordinate system employed in the tunnel was integrated in 1956 with the main baseline system south of the ice cliff, but the connecting traverse between the two systems could not be re-established in 1957. The 1956 coordinates given in Table I supersedes those presented by Goldthwait*.

MOTION ANALYSIS

The resultant deformation vector was in general directed principally forward (south), slightly downward, and inward to the void. The rate of forward motion increased rapidly upward from the base roughly following a parabolic curve tangent to the floor. Flow of the ice into the tunnel was essentially uniform over the span of the walls with a sharp decrease just above the floor. The downward component of movement was relatively small but erratic over the height of the wall and also decreased abruptly just above the floor. Thus, the basal few centimeters of ice blanketing the subglacial surface had virtually no motion. This fact coincides with observation that lichen on the boulder tops had remained undisturbed by the mobile ice mass. All evidence indicates that deformation proceeded at a uniform rate, and no erratic or backward motions, discontinuities, or significant cracks in the ice were detected.

Forward motion

The forward component of deformation in the tunnel resulted from the longitudinal compressive stress of the advancing ice cap (Fig. 8). Forward motion figures (Table II) were resolved in the actual direction of ice flow as hydrostatic closure deflected longitudinal strain vectors in toward the center of the tunnel. Eliminating the factor of ice creep into the tunnel, absolute forward motion was reduced to a direction approximately 4°W (clockwise) of the +Y' axis and from 4 to 10°E (counterclockwise) of the tunnel axis.

The vertical pattern of forward motion in the absolute plane roughly followed a parabolic curve with the greatest differential strain rates decreasing progressively upward from the base of the glacier (Table III). Movement proceeded in this pattern along each ring on both walls with pegs at comparable height above the floor having the same order of magnitude. A similar type of basal motion has been observed in the deformation of the pit at the inner end of the 1953 tunnel.

Shear strain in the basal few centimeters of the glacier was extremely small as it was only 0.03 mm/m-day at 0.05 m above the bouldery floor. The differential strain over a vertical section rapidly increased to 1.9 mm/m-day at 0.15 m above the floor, and then steadily decreased to 1.0 mm/m-day at 4.40 m above, to 0.5 mm/m-day at 30 m above, and to 0.3 mm/m-day at 40 m above (last two values are cliff face measurements). Thus, on the order to 2% of the forward motion occurred in the lowest 0.4% of the glacier height and 30% forward motion occurred at 10% height. Therefore, most of the gross forward motion near the terminus of this glacier resulted from a shearing action in the lowest portion of the ice mass.

A lateral comparison of forward motion rates at a common height along each wall showed in general less than 5% variation, except at the lower levels. The maximum discrepancy in the "C" pegs was 60%, but in the "B" pegs, which were most irregular, differences reached up to five times the lowest value. The range of values at comparable levels was more divergent along the west wall, largely because ring 4 values were consistently low. Dirt content in the ice at ring 4 was probably a partial retarding influence in the rate of forward motion.

* Goldthwait, R. P., Study of ice cliff in Nunatarssuaq, Greenland, U. S. Army Snow Ice and Permafrost Research Establishment, Corps of Engineers, Technical Report 39, 105p., in preparation.

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Table I. Peg and triangulation station coordinates.

Triangulation stations			
	X'	Y'	Z'
F1	10.000	10.000	10.000
F2	10.000	11.208	10.005
F3	10.056	12.345	10.109
F4	10.634	15.794	10.243
F5	11.228	19.866	10.484
F6	11.675	23.959	10.704

1956			Peg coordinates			1957		
	X'	Y'	Z'		X'	Y'	Z'	
Back face				Back face				
B	9.722	9.119	9.999	B	9.771	9.549	10.012	
C	9.741	9.245	10.342	C	9.804	9.826	10.337	
D	9.715	9.075	10.929	D	9.763	9.694	10.869	
E	9.686	8.841	11.548	E	9.785	9.986	11.472	
F	9.719	9.116	11.967	F	9.727	9.645	11.881	
UW	10.163	9.129	11.917	UW	10.183	9.801	11.790	
LW	10.164	9.131	11.182	LW	10.196	9.674	11.100	
UE	9.279	9.260	11.862	UE	9.354	10.046	11.727	
LE	9.282	9.263	11.246	LE	9.345	9.892	11.104	

Ring 1			Ring 1				
AE	9.121	10.139	9.905	AE	9.133	10.160	9.903
BE	9.048	10.153	10.067	BE	9.132	10.268	10.067
CE	9.017	10.166	10.353	CE	9.136	10.385	-
C'W	10.638	9.936	10.600	C'W	10.531	10.190	-
CW	10.644	9.895	10.137	CW	10.547	10.114	-
B'W	10.634	9.893	10.294	B'W	10.554	10.059	-
BW	10.614	9.897	10.144	BW	10.562	9.991	-
AW	10.571	9.893	10.004	AW	10.567	9.890	-

Ring 2			Ring 2				
AE	9.677	14.080	10.116	AE	9.689	14.074	10.105
BE	9.574	14.102	10.261	BE	9.637	14.165	10.271
CE	9.458	14.128	10.542	CE	9.615	14.339	10.542
DE	9.497	14.100	11.128	DE	9.775	14.590	11.111
EE	9.428	14.166	12.337	EE	9.690	14.910	12.232
EW	11.320	13.770	12.436	EW	11.142	14.569	12.334
DW	11.314	13.802	11.224	DW	11.090	14.338	11.154
CW	11.345	13.801	10.614	CW	11.193	14.100	10.567
AW	11.187	13.794	10.161	AW	11.179	13.803	10.155
LPW	11.346	13.894	11.428	LPW	11.139	14.489	11.371

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Table I (cont'd.)

1956			1957		
Peg coordinates			Peg coordinates		
X'	Y'	Z'	X'	Y'	Z'
Ring 3			Ring 3		
AE 10.091	18.064	10.347	AE 10.095	18.066	10.339
BE 10.070	18.077	10.493	BE 10.113	18.091	10.462
CE 10.061	18.080	10.756	CE 10.220	18.307	10.749
DE 10.069	18.098	11.394	DE 10.312	18.576	11.355
EE 10.046	18.105	12.386	EE 10.356	18.837	12.523
EW 11.609	17.732	12.576	EW 11.381	18.575	12.548
DW 11.518	17.825	11.430	DW 11.314	18.355	11.374
CW 11.493	17.813	10.837	CW 11.363	18.105	10.773
AW 11.447	17.807	10.389	AW 11.445	17.808	10.381
Ring 4			Ring 4		
AE 10.988	21.946	10.537	AE 10.994	21.947	10.538
BE 10.959	21.952	10.682	BE 11.080	22.063	10.629
CE 10.898	21.958	10.970	CE 11.071	22.165	10.916
DE 10.884	21.974	11.575	DE 11.142	22.463	11.500
EE 10.944	21.988	12.771	EE 11.260	22.772	12.637
FE 10.891	22.011	13.965	FE 11.199	22.936	13.859
YE 11.142	21.962	14.970	YE 11.325	23.230	14.764
YW 12.143	21.756	14.915	YW 12.160	23.040	14.790
HW 12.268	21.725	13.929	HW 12.095	22.654	13.883
G'W 12.250	21.720	13.253	G'W 12.071	22.580	13.222
GW 12.196	21.742	13.064	GW 12.011	22.558	13.035
FW 12.228	21.734	12.742	FW 12.037	22.503	12.715
E'W 12.223	21.724	12.428	E'W 12.029	22.465	12.404
EW 12.240	21.723	12.282	EW 12.050	22.435	12.256
D'W 12.235	21.703	11.614	D'W 12.058	22.156	11.580
DW 12.234	21.702	11.551	DW 12.058	22.148	11.517
CW 12.202	21.697	10.959	CW 12.059	21.883	10.922
BW 12.161	21.695	10.651	BW 12.107	21.776	10.608
AW 12.119	21.708	10.508	AW 12.117	21.705	10.512

Comparison of the over-all absolute motion on each wall shows significant anomalous values at certain lower pegs where small differences became critical (Table II). Differences under about 10% may have resulted from slightly unequal heights, again most prominent down low. Other variations may have been caused by irregularities in bottom topography or, as in ring 4, by dirt layers immediately below pegs.

Closure into tunnel 1

Movement of the tunnel boundaries inward resulted from hydrostatic pressure and longitudinal compressive stress of the moving ice mass applied at an oblique angle to the tunnel axis (Fig. 1-2, 9). Discussion of closure as such must necessarily be limited only to creep rate produced by the load of the overlying ice. Thus, closure figures are given in relation to the distances between the pegs for a common level on the same ring to eliminate the upward increasing effects of forward motion.

The effect of forward motion, causing appreciable increases of closure on the west side, can readily be seen in Figure 9. In comparison to the east side, the west side moved inward at an average of 30% faster at 0.15 m above the floor, 60% at 1.05 m, 120% at 2.25 m, and 230% at 3.40 m.

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Table II. Absolute forward motion in a plane 4°E of +Y' axis.

	Peg	Motion (mm)	Peg	Motion (mm)	% increase	Wall with higher value
Ring 1	AE	20	AW	3		
	BE	120	BW	90	33	East
			B'W	160		
	CE	230	CW	210	10	East
			C'W	250		
Ring 2	AE	*-6	AW	7		
	BE	60	BW	-		
	CE	220	CW	290	31	West
	DE	510	DW	510	-	-
	EE	760	EW	790	4	West
Ring 3	AE	2	AW	1		
	BE	20	BW	-		
	CE	230	CW	280	22	West
	DE	500	DW	510	2	West
	EE	760	EW	820	8	West
Ring 4	AE	1	AW	*-3		
	BE	120	BW	80	50	East
	CE	210	CW	180	17	East
	DE	510	DW	430	19	East
	EE	800	FW	750	7	East
	FE	940	HW	920	2	East
	YE	1280	YW	1290	1	East

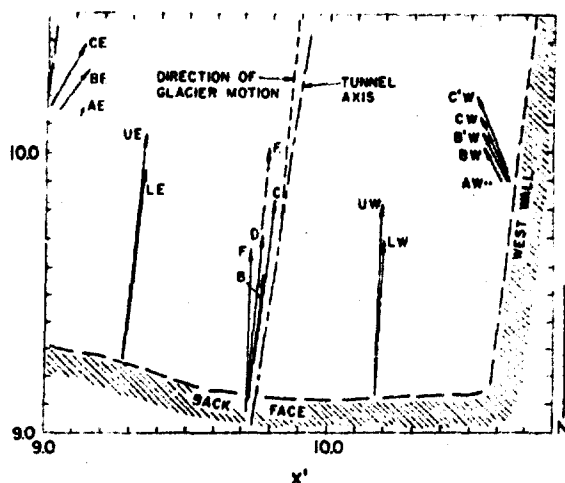
* Minute backward motion included as margin of error.

Table III. Absolute forward motion averages.

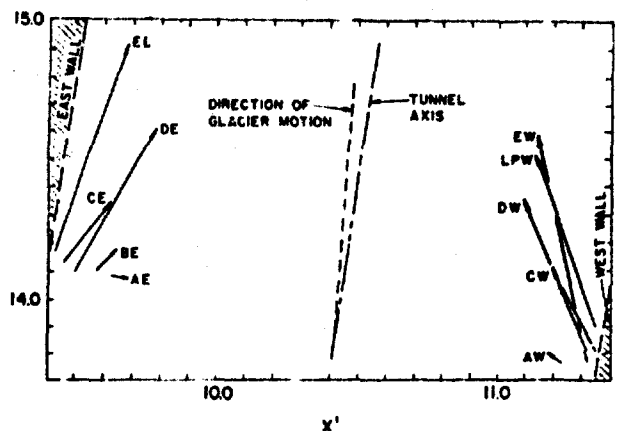
Approximate distance (m) above floor	Avg peg	Motion (mm)	Rate (mm/day)	Avg peg	Motion (mm)	Rate (mm/day)	Avg both walls (mm/day)	Differential strain rate above floor (mm/m-day)
0.05	AE	4	0.01	AW	4	0.01	0.01	0.3
0.15	BE	80	0.27	BW	85	0.29	0.28	1.9
0.45	CE	220	0.75	CW	240	0.82	0.79	1.8
1.05	DE	510	1.73	DW	480	1.63	1.68	1.6
2.25	EE	770	2.62	EW	790	2.69	2.66	1.2
3.40	FE	940	3.20	HW	920	3.13	3.17	0.9
4.40	YE	1280	4.35	YW	1290	4.39	4.37	1.0

Table IV. Horizontal closure, ring 4, 1956 tunnel.

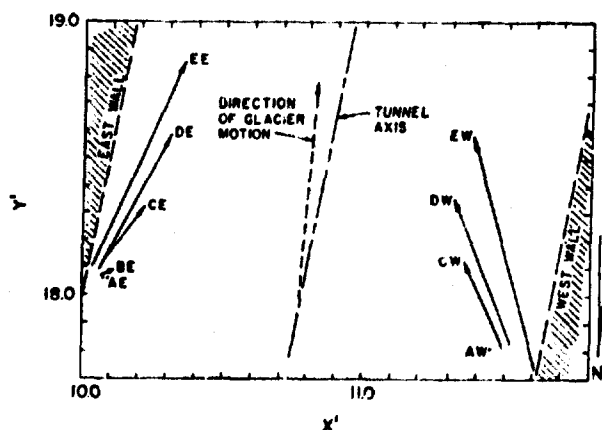
	Dist. above floor (m)	Mid-August 1956 dist. (cm)	Early June 1957 dist. (cm)	Difference (cm)	Rate 294 days (mm/day)
AE - BW	0.05	116	115	1	< 0.01
BE - BW	0.15	124	107	17	0.58
CE - CW	0.45	134	106	28	0.95
DE - DW	1.05	138	96	42	1.43
EE - FW	2.25	131	82	49	1.67
FE - HW	3.40	141	94	47	1.60



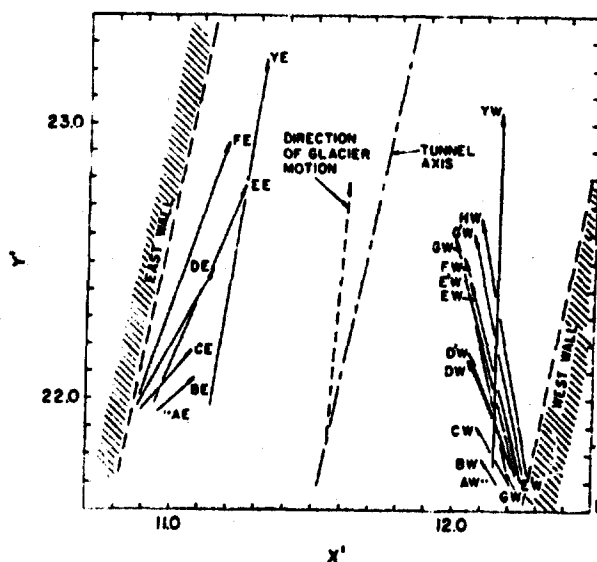
a. Back face and ring 1.



b. Ring 2.



c. Ring 3.



d. Ring 4.

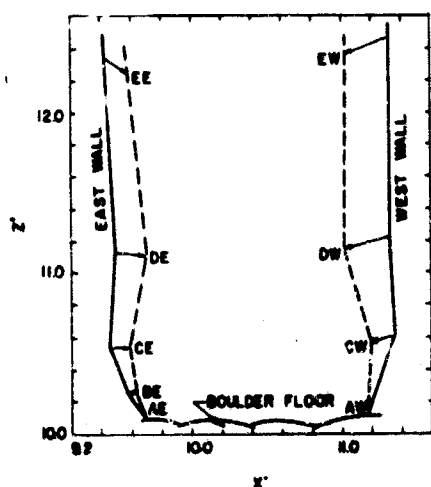
Figure 8. 1956 and 1957 X' and Y' peg coordinates.

The actual rate of approach between pegs of equal height rapidly increased upward to about 1 m above the floor and remained more or less constant above this point (Table IV). Closure at 0.05 m above the floor was almost negligible, less than 0.01 mm/day. At 0.15 m above the floor, closure had greatly increased to 0.58 mm/day and in the next upward 0.30 m it almost doubled to a rate of 0.95 mm/day. The closure rate gradually leveled off so that above 1 m the order of magnitude of closure was under 15% variation.

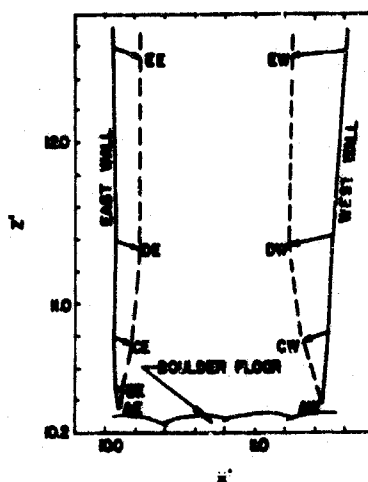
Downward motion

The resultant motion vector was directed slightly downward, although the north dipping ice foliation leads to the suggestion that net ice motion should have been at a gentle upward angle. The component of slump of the walls resulted from gravitational

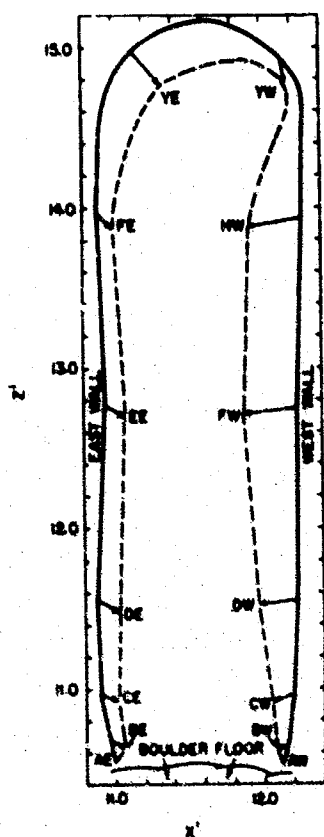
DEFORMATION OF RED ROCK ICE TUNNEL



a. Ring 2.



b. Ring 3.



c. Ring 4.

Figure 9. 1956 X' and Z' peg coordinates and closure vectors normal to tunnel axis.

force acting on the unsupported surface. This factor probably increased with time as an overhang developed above the relatively immobile ice at the glacier base.

The downward component of motion was very erratic both in lateral and vertical relationships (Table V). At 0.05 m above the floor, the downward motion was less than 10 mm. At 0.15 m above, slumping increased rapidly to a maximum of 50 mm. A significant correlation between figures of the upper portion of the wall can not be established other than to say they were generally higher than in the lower portion with a maximum net downward component of 110 mm.

The downward component of the ceiling arch as recorded on the "Y" pegs of ring 4 was from 130 to 210 mm and resulted from a combined effect of hydrostatic closure and slump.

Effect of dirt bands on motion

Several examples provided evidence that the motion of ice immediately overlying dirt layers was retarded (Fig. 3-5). Peg D'W of ring 4 moved forward slightly more slowly in relation to the differential motion of pegs both above and below (Fig. 7, 8d). In a sequence of pegs shown in Figure 10, pegs in the ice adjacent to

Table V. Downward component closure.

	Peg	Amount (mm)	Peg	Amount (mm)
Ring 1	AE	0.002		
	BE	0.000		
Ring 2	AE	0.011	AW	0.006
	BE	0.010 *	BW	-
	CE	0.000	CW	0.047
	DE	0.017	DW	0.070
	EE	0.105	EW	0.102
Ring 3	AE	0.008	AW	0.008
	BE	0.031		
	CE	0.037	CW	0.064
	DE	0.039	DW	0.056
	EE	0.063	EW	0.028
Ring 4	AE	0.001 *	AW	0.004 *
	BE	0.053	BW	0.043
	CE	0.054	CW	0.037
	DE	0.075	DW	0.034
	EE	0.084	FW	0.027
	FE	0.106	HW	0.046
	YE	0.206	YW	0.125

* Upward movement.

upper contact of the dirt bands consistently lagged in forward motion. Pegs BW and CW of ring 4 situated directly above dirt layers also moved comparatively slowly.

Core deformation

Cross-sectional shape changes of ice cylinders placed in the west wall of the tunnel were observed. The 10-cm diam cores placed in a hole of the same size in August, 1956, were distorted from the originally circular cross-section to an elliptical cross-section with a corresponding southward tilt (Fig. 11).

Partitions spaced at 90° and composed of cardboard or aluminum foil were set into three of the cores while they were being frozen. The cores were then emplaced in the ice walls with the partitions in essentially horizontal and vertical positions. The vertical dividers tilted southward on the order of 20° while the horizontal partition tilted from 4° southward to 8° northward. A tilt of 34° southward was recorded on a vertical partition of one core but inasmuch as the partition had separated, this value probably was invalid as true tilt.

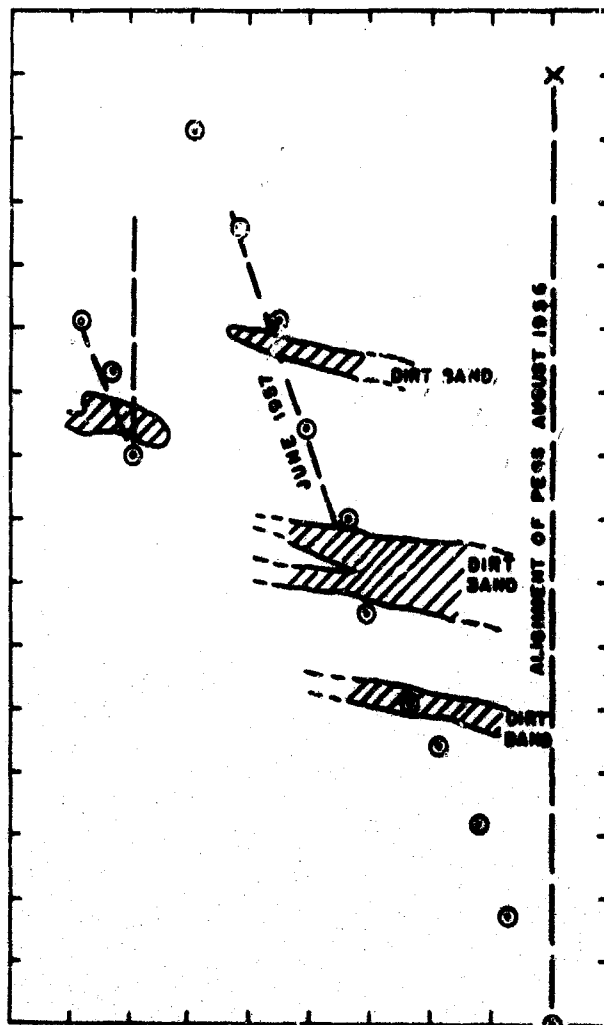


Figure 10. Showing tilt and lag of pegs above dirt bands. 1.7 m south of ring 4 on west wall (Fig. 7).

DEFORMATION OF RED ROCK ICE TUNNEL

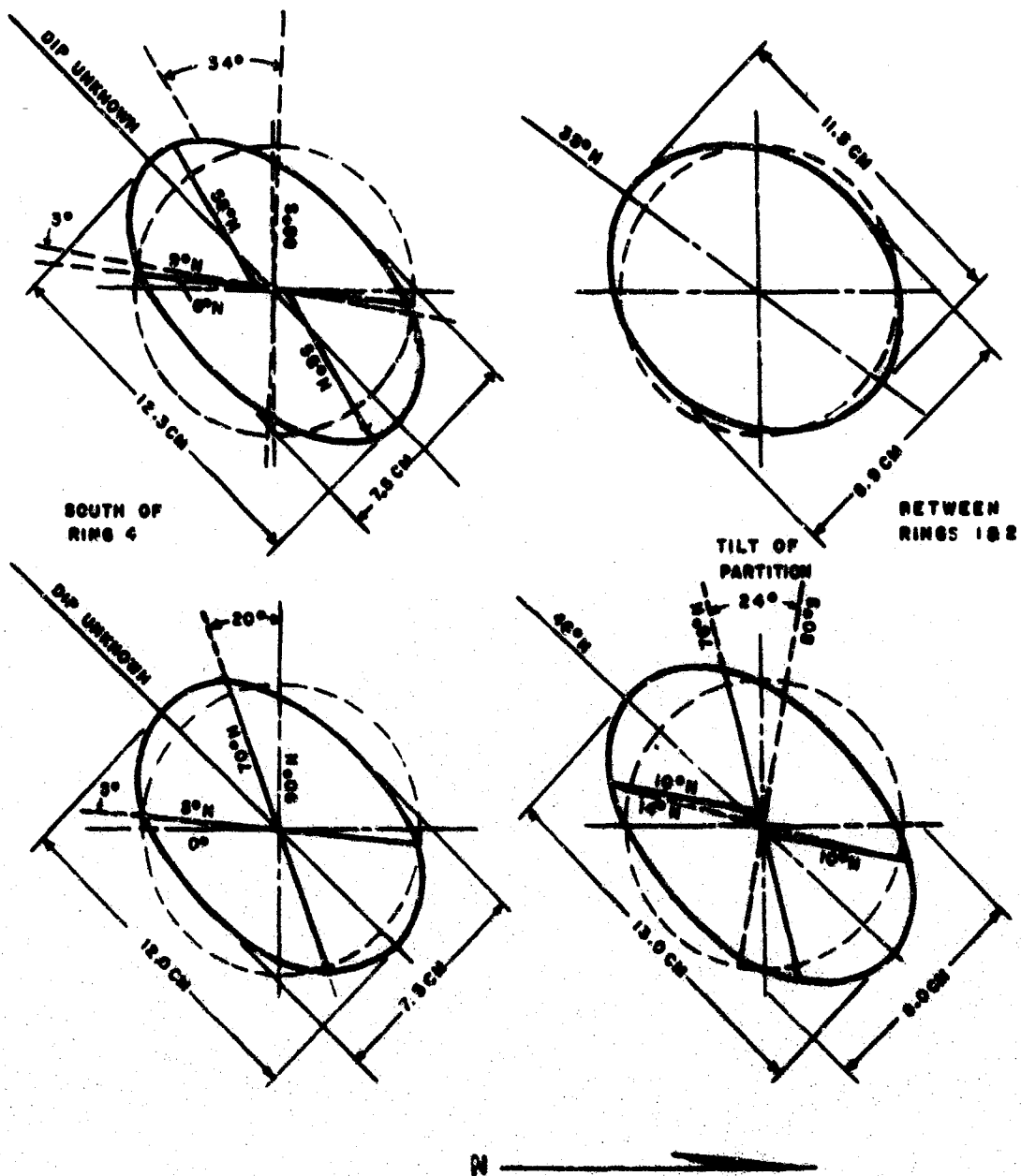


Figure 11. Deformation of cores and tilt of partitions, west wall of tunnel.